



Introductory guidance to atmospheric correction for use over coastal and inland waters

The signal measured by a satellite sensor is the sum of the three components (Fig. 1): 1/ Contribution from the atmosphere (including the scattering of sunlight by air molecules (Rayleigh scattering), aerosols, and the coupling between air molecules and aerosols); 2/ Contribution from the air-water interface (from foams and the reflection of the Sun and skylight on the water's surface); 3/ Contribution of the water body. The purpose of the atmospheric correction process is to remove the contribution of the atmosphere to the signal measured by the remote sensor, leading to the estimation of the radiance emanating from a water body (i.e., inland, coastal, or open ocean environment). The output of the atmospheric correction processing is the remote sensing reflectance, i.e. the ratio of the water-leaving radiance over downwelling plane irradiance. This signal can then be used to derive water bio-optical properties (e.g., total absorption and backscattering coefficients), biogeochemical variables (e.g., chlorophyll-*a*, sediments, and dissolved organic carbon concentration), as well as measures of transparency (see the [introductory](#) and [advanced](#) user's guides on bio-optical algorithms. Thus, the accuracy of the atmospheric correction step directly impacts the accuracy of any subsequent data processing steps. Some contributions (from air molecules, foams, reflection of the Sun and sky, atmospheric gases attenuation) can be estimated using ancillary data and viewing geometry (Mobley et al., 2016).

This document provides an introductory overview of atmospheric correction over coastal and inland waters. More in-depth guidance on advantages and limitations of each algorithm's type, intended for a more experienced audience, can be found in our [companion document](#). Other issues that occur over coastal and inland waters are addressed in the same document: 1/ absorbing aerosols and trace gases (urban pollution, dust) and 2/ adjacency effects.

Over open ocean waters, the contribution of the ocean to the total signal can be considered negligible (black pixel assumption) in the near-infrared (NIR) bands so the total signal is equal to the contribution of the atmosphere. Using the NIR bands allows aerosol models and optical properties to be estimated; these can in turn be used to estimate the atmospheric contribution to the measured signal in the visible bands. Once the contribution of the atmosphere is known, the estimation of the signal from the ocean is straightforward.

The black pixel assumption is not valid in more optically-complex waters (coastal and inland waters), where backscattering from suspended particles or from the seabed in shallow waters contributes to non-zero water reflectance in the red and NIR wavelengths (IOCCG, 2010). To overcome this challenge, many atmospheric correction algorithms were developed for ocean color remote sensors (SeaWiFS, MODIS-Aqua, MERIS, VIIRS, OLCI/S3/A/B) and high-spatial resolution sensors (OLI/L8 and MSI/S2A/B).

They can be grouped into seven different categories:

- (1) assumption of homogeneity of atmospheric and oceanic properties over a region of interest (De Keukelaere et al., 2018; Hu et al., 2000; Ruddick et al., 2000, 2006; Sterckx et al., 2015; Vanhellemont and Ruddick, 2021; Wang et al., 2021)
- (2) use of the shortwave infrared bands (Chen et al., 2014; He and Chen, 2014; Wang and Shi, 2005; Wang and Shi, 2007; Wang, 2007; Shi and Wang, 2009)
- (3) use of blue or ultra-violet bands (Oo et al., 2008; He et al., 2012; Wang and Jiang, 2018)
- (4) correction or modeling of the non/negligible signal from the ocean in the NIR (Moore et al., 1999; Siegel et al., 2000; Stumpf et al., 2003; Lavender et al., 2005; Bailey et al., 2010)
- (5) coupled ocean/atmosphere inversion based on artificial neural networks (Doerffer and Schiller, 2007; Schroeder et al., 2007; Fan et al., 2017) or optimization techniques (Brajard et al., 2006, 2012; Chomko and Gordon, 1998, 2001; Jamet et al., 2004; Kuchinke et al., 2009; Stamnes et al., 2003; Steinmetz et al., 2011)

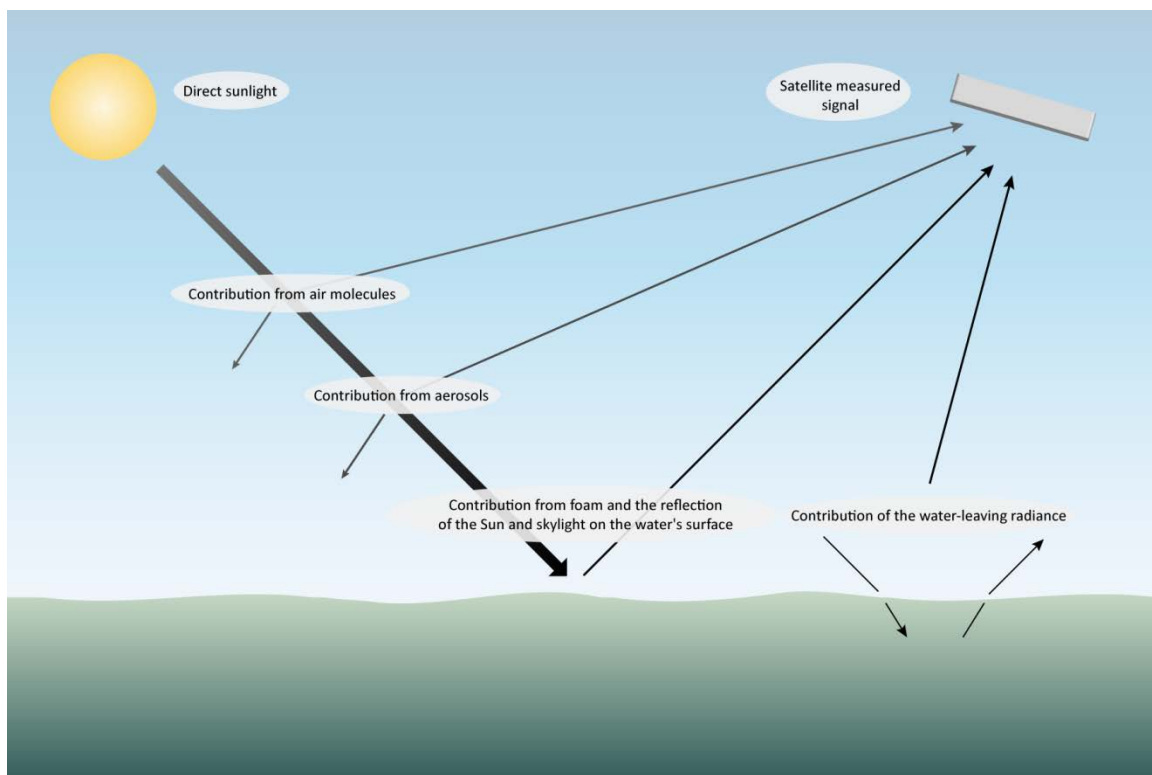


Figure 1. Schematic path of the sunlight through the atmosphere and the water body as seen by a space-borne sensor.



(6) based on statistical relationships/hypothesis to estimate the aerosol and water components (Gossn et al., 2019, 2021; He and Chen, 2014; Ibrahim et al., 2019; Saulquin et al., 2016; Singh and Shanmugam, 2014).

(7) based on observations from multiple observation angles (Xu et al., 2016; Wang et al. 2020)

Validations through match-up exercises have been provided for ocean color sensors (Jamet et al., 2011; Goyens et al., 2013; IOCCG report #10 and in preparation) and high-spatial resolution sensors (Pahlevan et al., 2021; Xu et al., 2020).

Authorship information

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Introductory Atmospheric Correction Feedback to: Mbneely@geoaquawatch.org.

Date

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